## Arctic Change 2014



# SYNTHESIS OF PERMAFROST RESEARCH ALONG THE EURASIA ARCTIC TRANSECT

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(Bely Island) -137.4

Tambey

38/0.33

# ECI DADA COLLECTION

- 1. General physiographical and climatic information for the transect.
- 2. Summary of the existing information on vegetation, soils and permafrost for the transect locations and Yamal as a whole.
- 3. Summary of the own data on vegetation, soils and permafrost.
- 4. PhD dissertations of the participants devoted to vegetation and permafrost of Central Yamal, partly translated from Russian into English.
- 5. List of publications and presentations using results of LCLUC-Yamal project.
- 6. Maps cartographic models of vegetation-permafrost interrelation.

Permafrost research along the Eurasia Arctic Transect (EAT) included estimation of the thermal state of permafrost and activity of cryogenic processes through ground measurements and remote-sensing data analysis. Transect covers the distance of 1880 km from Nadym (65°18'51"N, 72°51'43"E) to FJL Hayes Island (80°35'56"N, 57°54'20"'E):

LOCATIONS	DISTANCE, km
Nadym - Laborovaya	340
Laborovaya – Vaskiny Dachi	290
Vaskiny Dachi – Kharasavey	125
Kharasavey – Bely Island	260
Bely Island – Hayes Island (Krenkel)	865

Yamal being an area of rapid natural and anthropogenic changes is an excellent object to calculate spatial distribution and temporal dynamics of ground temperature and active layer depth. Permafrost is affected by natural surface disturbances (cryogenic landslides, thermodenudation, thermoerosion, thermokarst), anthropogenic (structures, vehicle tracks, sandpits), and natural-anthropogenic (reindeer pasturing).

Factors specific for research polygons in Tundra zone with continuous permafrost distribution

Laborovaya: Piedmont of Urals. Deeply dissected surface, combination of well drained tops and slopes with poorly drained valleys and lake depressions. Sandy and clayey deposits with clasts and bedrock outcrops.

Vaskiny Dachi: Terraced Middle to Late Pleistocene plain. Deeply dissected surface, combination of convex drained sandy and concave poorly drained clayey tops and slopes, as well as poorly drained concave and well drained convex lake depressions and valleys. Peat on concave tops and valleys.

**Kharasavey:** Terraced Late Pleistocene plain. Moderately dissected, poorly drained both sandy and clayey deposits on tops and slopes and peat in the valleys.

Bely island: Terraced Holocene plain. Not deeply dissected, flat, gentle slopes, poorly drained. Drained sandy top edges. Peat in the valleys.

Hayes island: Terraced Holocene and Late Pleistocene plain with bedrock dykes forming long narrow ridges. Deeply dissected, terraced slopes, poorly drained except dyke tops and steep sandy slopes. Limited organic matter on all surfaces. In discontinuous permafrost zone, in northern taiga forcing factors are different. Permafrost as a feature (thus Tg and ALD) appear only away from forests, at the mires and peat plateaus.

Nadym: Terraced Holocene and Late Pleistocene plain. Relatively flat, subhorizontal, poorly drained except for the frost-heave mounds. Peat on all surfaces.

**Krenkel** location

Krenkel-2

**Krenkel-1** 

osef L/and

Hayes Islan

Barents S

Average air

temperature

-13,3°C

-10,3°C

-9,7°C

-8,1°C

-7,0°C

Quaternary

geology

Highly variable topography, deep dissection in the central part of Yamal compared to relatively flat southern and northern parts of EAT determine the role of precipitation, both winter and summer. Main forcing factor for ground temperature is snow cover thickness. It is the highest on slopes and in the narrow valleys. Snow insulates surface not only from winter cold, but also from summer warming, because snow patches survive at least till mid-July thus reducing active layer depth. At the same time, at the hilltops snow is blown away along with vegetation cover resulting in the lowest ground temperature, but at the same time, deepest summer thaw.



### Thaw depth





### LIST OF ECI LCLUC-RELATED AND OTHER YAMAL PUBLICATIONS

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Ground temperature at Bely Ostrov the coldest in sandy windblown terrace (OB3), medium in loamy highly vegetated slope (OB2) and warmest in the stream valley with snow accumulation (OB4)

Surface disturbances of all kind as a rule increase snow cover thus increasing ground temperature, and at the same time increase active layer depth because vegetation cover is reduced. Landslides and thermocirques specifically for Yamal expose saline permafrost with complicated phase transition process resulting in seasonal thaw different from the depth of zero temperature. In addition, forcing factors interact with each other. Spatial pattern of vegetation cover is mainly determined by landslide and



### Conclusion

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Ă 60

100

thermodenudation activity the last several millennia thus affecting distribution of the active layer depth.



Snow cover distribution, results of modelling based on field snow survey(left), and landscape-based (right)

In Central Yamal measured ground temperature varies from -0.3°C up to -7°C, active layer depth varies from 40 to 240 cm. Temporal variations caused by climate fluctuations range at  $\pm 2^{\circ}$ C for ground temperature, and  $\pm 5-10$  cm for active layer depth. Thus, spatial factors, such as lithology and surface covers are of much higher importance compared to climatic factors.

Spatial distribution of permafrost parameters along the EAT involves latitudinal zonality based on directional lowering of air temperature northward from Nadym site in the south to Hayes site in the north. It is established that, on the whole, consistent trend of bioclimatic subzones northward determines the consecutive change of various parameters of permafrost. However, local factors connected to relief, drainage degree, location of plots on different landforms, which determine snow accumulation and vegetation mat thickness, distort zonal pattern.

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